# Assessment of the Critical Load of Trace Elements in Soils Polluted by Pyrite tailings. A Laboratory Experiment

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Abstract Nineteen soil samples (SE Spain) with very different chemical physical properties and developed over different parent materials were contaminated by adding increments of an acidic solution from oxidised pyrite tailings. The quantities of Cu, Zn, Cd and Pb precipitated by the soil samples were directly and significantly related to the pH-buffering capacity. However, when the contamination caused the pH to fall below 3.0 the soil tended to release a fraction of the element adsorbed, which increased as the pH decreased. The quantity of each precipitated element at which the action value for each element is reached, was also directly related to the pH-buffering capacity. Nevertheless, in carbonate-rich soils, the precipitated Cu and Zn maintained a relatively high level of bioavailability, while Cd reached a critical level with a content exceeding 25 mg kg<sup>-1</sup>, regardless of the pHbuffering capacity.

**Keywords** Soils · Pyrite tailings · Trace elements · Buffering capacity · Critical Load · ph

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## **1** Introduction

Under natural conditions, the concentration of trace elements in the soil is relatively low, but certain human actions, such as agriculture, (Kabata-Pendias 1995), the application of tailings with residues when used as an organic amendment (Plaquart et al. 1999; McBride 2003), the application of solid urban waste as fertiliser, atmospheric particle transport (Alloway 1995; Föstner 1995), as well as mining (Egger 1994; Ripley et al. 1996; Vartanyan 1989), mining accidents (Cabrera et al. 1999; Simón et al. 1998) or continuous atmospheric depositions (Nicholson et al. 2003) can raise contamination above toxic levels. In these cases, the soil is capable of attenuating the mobility of diverse contaminants by, for example, filtration, neutralisation, adsorption and precipitation; these processes inactivate potential toxicity, blocking undesired elements from passing to more sensitive systems such as the air or water. The capacity of each soil to retain metals depends on the soil properties, mainly texture, organic-matter content, ion-exchange capacity, oxide contents, pH, specific surface area and carbonate content (Alloway 1995; Ross 1994). When this capacity is exceeded, the soil ceases to be effective as the protector of the ecosystem (sink) and can even function as a source of toxic substances. The critical load represents "the maximum quantity of a given contaminant that can be supplied to a soil without causing chemical changes leading to long-term harmful effects on ecosystem structure and function" (Hettelingh et al. 1991).

This study was aimed at estimating the critical load of soils with different properties withstanding a contaminating solution resulting from the oxidation of pyrite tailings with a pH value lower than 2.0 and containing a high concentration in trace elements such as Cu, Zn, Cd and Pb. This estimation can be made as a function of: (a) maximum concentration of Cu, Zn, Cd and Pb that precipitate in each soil within the range of pollution considered in this study (MPC, maximum precipitated concentration), (b) concentration of precipitated elements at which action values (AV, action values) or values above which plants and soil solution should be watched (Prüeß 1997), are reached (CAV).

When the soil concentration of a certain contaminant approaches the MPC, not only does it exceed its critical load and affect the functioning of the ecosystem, but new additions could reach surface waters and groundwater whereupon the contamination would spread to other ecosystems. At the same time, given that the precipitated elements may remain in bioavailable form, before the MPC is reached, the AV may be reached, at which point the soil would be expected to begin to display harmful effects. Therefore, the structure and functioning of the ecosystem would begin to be visibly affected when the soil reaches the AV, whereas when it reaches the MPC the surrounding ecosystems would also be visibly affected.

### 2 Materials and Methods

### 2.1 Soil Properties

The soil samples belong to the provinces of Granada and Jaen located in the southeast of Spain. The selected soil samples represent a heterogeneous range of soils with different chemical properties and developed over different parent materials (Table 1).

Soil samples were air dried and then screened to 2 mm to analyse them. Particle size distribution was determined by the pipette method after elimination of organic matter with  $H_2O_2$  and dispersion with sodium hexametaphosphate (Loveland and Whalley 1991). The pH was measured potentiometrically in a 1:2.5 soil–water and soil 0.1 mol 1<sup>-1</sup> KCl suspension. The calcium carbonate equivalent content (CaCO<sub>3</sub>, %) was determined by the method of Bascomb (1961). Total carbon was analysed by dry combustion with a LECO SC-144DR instrument. Organic carbon (OC)

was determined by the difference between total carbon and inorganic carbon from CaCO<sub>3</sub>. The cation-exchange capacity (CEC) was determined with 1 M Na acetate at pH 8.2, measuring sodium in a Meteor NAK-II flame-photometer. Exchange bases were determined with 1 M H<sub>4</sub>N acetate at pH 7.0, measuring sodium and potassium in a Meteor NAK-II flame-photometer and calcium and magnesium by atomic absorption spectrometry in a Varian SpectrAA 220FS machine, the specific surface area (SA) determined by weighing the water adsorbed by the sample from a solution saturated with CaCl<sub>2</sub> (Keeling 1961); total iron (Fe<sub>d</sub>) and aluminium (Al<sub>d</sub>) oxides and oxi-hydroxides of the soil samples were extracted with citrate-dithionite (Holmgren 1967).

# 2.2 Contaminating Solution (CS)

Ten grams of pyrite tailings from the mining spill at Aznalcóllar (Seville, Spain) were placed in contact for 3 days with 1,000 cm<sup>3</sup> of  $H_2O_2$  at 15%. The contaminating solution (CS) presented a pH of 1.7 and a concentration in Cu, Pb, Zn and Cd of 14.5, 3.2, 55.4 and 0.21 mg dm<sup>-3</sup>, respectively (more details of the contaminating solution composition is given in Simón et. al. 1999).

# 2.3 Spiked Soils

According to their properties, different quantities of CS were added to soil samples. Thus, 1, 1.5, 2, 3 and 5 cm<sup>3</sup> CS g<sup>-1</sup> were added to the sandy and acid soils (pH<7.0); 1, 2, 4, 6 and 10 cm<sup>3</sup> CS g<sup>-1</sup> were added to soils with a sandy-loam texture or finer and pH> 7.0 and CaCO<sub>3</sub> <2%; 2, 6, 10, 15 and 20 cm<sup>3</sup> g<sup>-1</sup> were added to soils with a loamy texture or finer, pH>7.5 and CaCO<sub>3</sub> between 2% and 30%; and 10, 20, 30, 40 and 50 cm<sup>3</sup> g<sup>-1</sup> were added to soils with a loamy texture or finer, pH>8.0 and CaCO<sub>3</sub>>30%. After 72 h of shaking (Alonso et al. 1997), each soil extract–CS was measured for pH and centrifuged at 3,000 rpm for 15 min, separating the solid and liquid fractions.

# 2.4 Total Water and NH<sub>4</sub>NO<sub>3</sub> Extractable Trace Element Concentrations

To determine water soluble forms, soil-water extracts were prepared in a ratio 1:10 (Norma DIN 38 414-4)

Table 1	Physico-chemical	properties	of the	studied	soils
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Samples	Parent material	pH <sub>H2O</sub>	рН <sub>КС1</sub>	CaCO <sub>3</sub> (%)	M.O. (%)	cmol <sub>c</sub> kg <sup>-1</sup>			V (%)
						Ca	Mg	CEC	
J-1	Quartzite	6.90	5.90	0.20	1.62	2.50	0.58	3.51	35.4
J-2	Sandstone	6.80	5.50	0.70	0.96	4.25	0.82	5.62	85.8
J-3	Granite	5.70	4.10	0.80	5.06	2.75	0.41	8.79	46.2
J-4	Slate	6.40	4.40	0.10	1.98	6.75	2.88	7.73	100
Gr-1	Filite	8.20	7.20	3.30	2.50	6.00	2.99	5.62	100
Gr-2	Conglomerate	6.60	4.70	1.00	1.01	10.87	3.67	16.17	77.8
Gr-3	Limestone	8.00	7.00	0.40	8.70	23.00	1.16	24.6	78.1
Gr-4	Schist	7.60	6.20	0.60	3.69	9.50	2.66	12.65	100
GR-5	Schist	7.70	6.60	0.70	1.93	6.50	1.08	7.03	96.5
Gr-6	Conglomerate	8.30	7.10	0.60	1.74	16.21	2.49	18.98	100
Gr-7	Limestone	7.80	6.90	0.90	7.90	28.00	1.49	30.93	74.7
Gr-8	Filite	8.30	7.50	19.60	5.77	12.25	5.06	15.42	83.9
Gr-9	Limestone	8.40	7.20	5.10	2.39	27.00	1.33	16.17	100
Gr-10	Evaporite	7.80	7.50	56.50	1.90	41.89	4.32	13.71	100
Gr-11	Limestone	8.20	7.40	22.90	2.71	30.50	1.91	18.98	100
Gr-12	Limestone	8.20	7.50	59.00	3.34	27.25	1.49	15.47	100
Gr-13	Marl-Limestone	8.40	7.20	79.60	0.85	24.25	1.00	13.36	100
Gr-14	Dolomite	8.20	7.50	32.60	3.08	10.00	3.24	11.25	100
Gr-15	Dolomite	8.30	7.30	84.60	1.90	4.25	3.15	12.30	100
Sample	% Sand	% Silt	% Clay	SE (m <sup>2</sup> /g)	Fe <sub>d</sub> %	$\frac{BC_{3.5}}{(cmolH^+ kg^{-1})}$			
J-1	59.30	28.00	12.70	4.54	0.34	0.62			
J-2	66.30	19.20	14.50	13.73	0.73	0.39			
J-3	78.00	14.00	8.00	17.66	0.26	0.47			
J-4	57.40	32.00	10.60	34.81	0.92	0.85			
Gr-1	58.80	23.30	17.90	14.03	0.87	8.36			
Gr-2	32.50	13.30	54.20	108.45	2.10	1.14			
Gr-3	60.80	16.10	23.10	69.27	0.76	8.28			
Gr-4	67.60	21.30	11.10	43.86	0.87	1.93			
Gr-5	68.10	19.20	12.70	19.45	1.55	2.91			
Gr-6	33.60	15.60	50.80	100.56	2.83	5.81			
Gr-7	23.60	55.80	20.60	111.88	2.49	16.14			
Gr-8	49.80	28.30	21.90	32.78	1.02	17.76			
Gr-9	41.20	27.30	31.50	76.51	1.84	20.07			
Gr-10	22.40	55.70	21.90	43.68	0.21	90.83			
Gr-11	43.60	31.60	24.80	78.47	0.84	79.52			
Gr-12	38.80	37.10	24.10	54.58	0.55	82.49			
Gr-13	22.40	41.00	36.60	62.80	0.26	81.79			
Gr-14	40.80	33.60	25.60	29.78	1.10	63.34			
Gr-15	38.30	58.10	3.60	3.58	0.21	50.31			

Soil samples are labelled according to the location of the sampling site (Jaen province, J and Granada province, Gr)

M.O. organic matter, Ca exchangeable calcium content, Mg exchangeable magnesium content, CEC cation-exchange capacity, V base saturation percentage, Ss Specific surface, Fed iron oxides

and then Cu, Pb, Zn and Cd water extractable concentrations were analysed. Trace element precipitated concentrations (MPC) were determined as the difference between trace element concentration in the CS and water extractable concentration measured in soil water extracts. Finally, the pH value measured in the soil–water extracts. Soils were also analysed for the trace element concentrations extractable by NH<sub>4</sub>NO<sub>3</sub> 1 M (DIN 19730; Prüeß 1997). In all cases, the concentration of the different metals (Cu, Zn, Cd and Pb) was measured by ICP-MS in a PE SCIEX ELAN-5000 A spectrometer.

Distribution coefficients ( $K_d$ ) represent the sorption affinity of the soil solid phase for trace elements in solution and can be used as a valuable tool to study metal-cation mobility and retention in soil systems. The distribution coefficients ( $K_d$ ), ratios between the quantity of precipitated element (mmol kg<sup>-1</sup>) and the quantity remaining in solution (mmol dm<sup>-3</sup>), were calculated according to Alloway (1995). According to Anderson and Christensen (1988), high  $K_d$  values indicate that the metal has been retained by the solid phase through sorption reactions, while low values of  $K_d$  indicate that a large fraction of the metal remains in solution.

# 2.5 Determination of Buffering Capacity of Soils

The pH-buffering capacity of soils at pH=3.5 was evaluated (BC<sub>3.5</sub>) as the cmol of  $H^+$  needed to reduce soil pH value up to 3.5. The titration curves were drawn following the method of Hartikainen (1986), soil samples (air dried, <2 mm) were equilibrated with solutions of increasing concentrations of 10<sup>-3</sup> M HCl (soil-solution ratio of 1:10) while maintaining the ionic strength constant. The suspensions were shaken for 60 min and then the pH measured after 4 days of settling; then the amount of H<sup>+</sup> added were plotted against the equilibrium pH values (titration curves). From these curves, the values for the acid neutralisation (BC) were determined graphically, this being referred to as the quantity of acid needed to bring the pH of the soil to a given value (Van Breemen et al. 1983). In the present study, the pH value of 3.5  $(BC_{3,5})$  was taken as the reference value, like one of the intervals defined by Ulrich (1981) for soils and reflects the aluminium and iron buffer region. Finally, the data were statistically processed using SPSS 12.0 software package.

# **3** Results and Discussion

# 3.1 pH-Buffering Capacity

The main property that marked a significantly different behaviour among the soils was the BC<sub>3.5</sub>, which was related (P<0.005) to the calcium carbonate content (% CaCO<sub>3</sub>) and the sum of the exchangeable calcium and magnesium content in soils ([Ca+Mg] expressed in cmol<sub>c</sub> kg<sup>-1</sup>):

BC<sub>3.5</sub>(cmol H<sup>+</sup> kg<sup>-1</sup>) = 0.792 CaCO<sub>3</sub> + 0.870[Ca + Mg]  
$$r^2 = 0.903$$

indicating that the weathering of the carbonates and the exchange between basic cations (among which Ca and Mg are the most abundant) and protons are the main mechanisms controlling soil acidification. In any case, soils with different BC<sub>3.5</sub> values show similar pH-buffering capacity pattern in such a way that, based on this pattern, the soils were assigned to three groups (Fig. 1):

- soils with BC<sub>3.5</sub><3 cmol H<sup>+</sup> kg<sup>-1</sup>, in which the pH value was reduced rapidly with the addition of CS, reaching pH values of 3.0 for additions of ≤2 cm<sup>3</sup> CS g<sup>-1</sup> of soil;
- soils with BC<sub>3.5</sub> between 3 and 20 cmol H<sup>+</sup> kg<sup>-1</sup>, in which the pH decreased gradually, reaching values of 3.0 for relatively lower concentration of contaminating solution (between 2 and 4 cm<sup>3</sup> CS g<sup>-1</sup> of soil);
- soils with BC<sub>3.5</sub>>20 cmol H<sup>+</sup> kg<sup>-1</sup>, in which the decline of the pH was very gradual and in no case reached values of less than 6.5.



**Fig. 1** Changes in pH values measured in H<sub>2</sub>O of soils with different pH-buffering capacity at pH=3.5 (BC<sub>3.5</sub>) with increasing amounts of contaminating solution added (cm<sup>3</sup> SC g<sup>-1</sup> soil). (*A*) BC<sub>3.5</sub>>20 cmol H<sup>+</sup> kg<sup>-1</sup>; (*B*) BC<sub>3.5</sub> between 3 and 20 cmolH<sup>+</sup> kg<sup>-1</sup>; (*C*) BC<sub>3.5</sub><3 cmol H<sup>+</sup> kg<sup>-1</sup>

# 3.2 Water-extractable Concentrations and Coefficient of Distribution

The coefficient of distribution ( $K_d$ ), the ratio between the quantity of the precipitate of each element (mmol kg<sup>-1</sup>) and the quantity that remained in solution (mmol dm<sup>-3</sup>), revealed that the precipitation rate of the different trace elements was related to the BC<sub>3.5</sub>. Thus, the greatest proportion of water extractable trace elements concentration was found in soils with BC<sub>3.5</sub><3 cmol H<sup>+</sup> kg<sup>-1</sup> (Fig. 2 A), followed by those with BC<sub>3.5</sub> between 3 and 20 cmol H<sup>+</sup> Kg<sup>-1</sup> (Fig. 2 B); meanwhile, the greatest precipitation rate was registered for soils with BC<sub>3.5</sub>> 20 cmol H<sup>+</sup> Kg<sup>-1</sup> (Fig. 2 C). Similarly, the  $K_d$  values showed high solubility of Cd and Zn as opposed to Cu and Pb.

In general, within each of the defined groups, the quantity of precipitated metal tended to increase progressively with the ratio CS–soil extract until the maximum retention capacity of each soil (MPC), above which the precipitated metal concentration either tended to remain constant or tended to diminish. The first behaviour was noted in the soils with BC<sub>3.5</sub>>20 cmol H<sup>+</sup> kg<sup>-1</sup> (Fig. 3, Gr13), whereas the soils with BC<sub>3.5</sub><20 cmol H<sup>+</sup> kg<sup>-1</sup> displayed the second behaviour (Fig. 3, Gr3). Therefore those soils with BC<sub>3.5</sub><20 cmol H<sup>+</sup> kg<sup>-1</sup> could thus become a source of contamination when the MPC is exceeded (chemical time bomb; Stigliani 1988). In general, this



**Fig. 2** Distribution coefficient for Cu, Pb, Cd and Zn in the studied soils, calculated at a dosage of 500 mg kg<sup>-1</sup> of SC ( $K_{d500}$ ). In the graph, we can distinguish between (*A*)  $K_{d500}$  of soils with BC<sub>3.5</sub><3 cmol H<sup>+</sup> kg<sup>-1</sup>; (*B*)  $K_{d500}$  of soils with BC<sub>3.5</sub>> between 3 and 20 cmolH<sup>+</sup> kg<sup>-1</sup>; (*C*)  $K_{d500}$  of soils with BC<sub>3.5</sub>> 20 cmol H<sup>+</sup> kg<sup>-1</sup>



Fig. 3 Graph showing the precipitated concentration of copper with the increasing amount of metal (copper) added with the contaminating solution in soils Gr-3 (with BC<sub>3.5</sub><20 cmol  $H^+$  kg<sup>-1</sup>) and Gr-13 (with BC<sub>3.5</sub>>20 cmol  $H^+$  kg<sup>-1</sup>)

process takes place in the studied soils when the pH of the soil extract-CS is less than or equal to 3.0.

In any case, these two types of behaviour should be understood within the range of the studied pollution, since at higher contamination rates even the pH of those soils with  $BC_{3.5}>20$  cmol  $H^+$  kg<sup>-1</sup> could fall to values below 3.0, releasing precipitated elements and thus becoming in a potential threat to the ecosystem exposed to them.

Although MPC<sub>Cu</sub>, MPC<sub>Pb</sub>, MPC<sub>Zn</sub> and MPC<sub>Cd</sub> were significantly related (p<0.01) to the BC<sub>3.5</sub>, this relationship presented certain particularities, depending on the element. Thus, in the case of Cu, Zn and Cd, the relationship between the two parameters was linear:

$$\begin{split} MPC_{Cu} &(mg\,kg^{-1}) = 11.84 + 3.11\,BC_{3.5}\,(cmol\,H^+\,kg^{-1})r^2 = 0.965\\ MPC_{Zn} &(mg\,kg^{-1}) = 13.03 + 9.35\,BC_{3.5}\,(cmol\,H^+\,kg^{-1})r^2 = 0.955\\ MPC_{Cd} &(mg\,kg^{-1}) = 0.124 + 0.02\,BC_{3.5}\,(cmol\,H^+\,kg^{-1})r^2 = 0.976 \end{split}$$

while Pb presented a different behaviour, depending on the value of the BC<sub>3.5</sub> (Fig. 4). When the BC<sub>3.5</sub> < 70 cmol H<sup>+</sup> kg<sup>-1</sup>, MPC<sub>Pb</sub> was related logarithmically to BC<sub>3.5</sub> and tended to reach a constant value. When





**Fig. 4** Relationship between the maximum quantity of Pb precipitated (MPC<sub>Pb</sub>) and the pH-buffering capacity (BC<sub>3.5</sub>)

 $BC_{3.5}$ >70 cmol H<sup>+</sup> kg<sup>-1</sup>, as in the soils with a very high CaCO<sub>3</sub> content, the MPC<sub>Pb</sub> tended to increase linearly.

3.3 Copper, Lead, Zinc and Cadmium NH<sub>4</sub>NO<sub>3</sub> Extractable Concentrations

The relationship between the precipitated and the NH<sub>4</sub>NO<sub>3</sub> 1 M extractable concentration for each metal was established. From this relationship, we can deduce the quantity of precipitated metal at which the action level is reached (CAV) accepted by the DIN 19730 guidelines where AV are reported as 1 mg Cu kg<sup>-1</sup> soil (AV<sub>Cu</sub>), 3 mg Pb kg<sup>-1</sup> soil (AV<sub>Pb</sub>), 5 mg Zn kg<sup>-1</sup>  $(AV_{Zn})$  and 0.08 mg Cd kg<sup>-1</sup> soil  $(AV_{Cd})$  (Prüeß 1997). In all cases, the CAV value was also significantly related (p < 0.01) to the BC<sub>3.5</sub>, although with certain particularities (Fig. 5). Thus, the CAV of Cu and Zn were related to the BC<sub>3.5</sub> through seconddegree equations, differentiating among soils with  $BC_{3.5} < 70 \text{ cmol H}^+ \text{kg}^{-1} \text{ and } BC_{3.5} > 70 \text{ cmol H}^+ \text{kg}^{-1}$ . It is worth noting that the soils with  $BC_{3,5}>70$  cmol  $H^+$  kg<sup>-1</sup> presented relatively low CAV values, indicating that, although the highly carbonate soils can precipitate high quantities of Cu and Zn, their bioavailability is high and can easily get into the food chain.

On the contrary, Cd and Pb did not differ. Thus, in all the soils, the  $CAV_{Cd}$  was related to the  $BC_{3.5}$ according to the logarithmic equation in which there was a tendency for a steady value to be reached (Fig. 5c); this indicates, in general terms, that when the soils presented a Cd concentration of around 2.5 mg kg<sup>-1</sup>, they exceeded the critical bioavailability level. In the case of Pb (Fig. 5d), the relation between  $CAV_{Pb}$  and  $BC_{3.5}$  was exponential, so that an increase in the neutralisation capacity of soil acids allows a considerable increase in Pb before the critical bioavailability level is reached.

#### **4** Conclusions



Fig. 5 Relationship between the quantity of metal precipitated (CAV<sub>Cu</sub> and CAV<sub>Zn</sub>) at which the action value (AV) is reached and the pH-buffering capacity at pH 3.5 (BC<sub>3.5</sub>). Relationship between the quantity of metal precipitated (CAV<sub>Cd</sub> and CAV<sub>Pb</sub>) at which the action value (AV) is reached and the pH-buffering capacity at pH 3.5 (BC<sub>3.5</sub>)

Cd and Pb in soils. Thus, while the  $NH_4NO_3$  extractable Cu and Zn concentration tends to increase in highly carbonated soils, the Pb reduces drastically and the Cd tends to remain constant regardless of the soil characteristics.

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